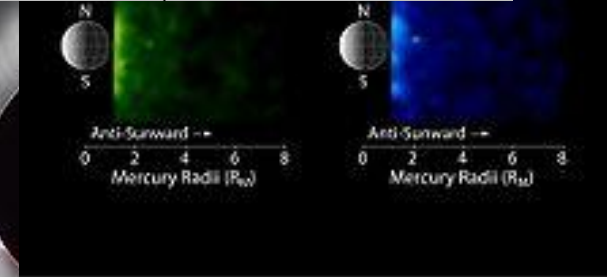
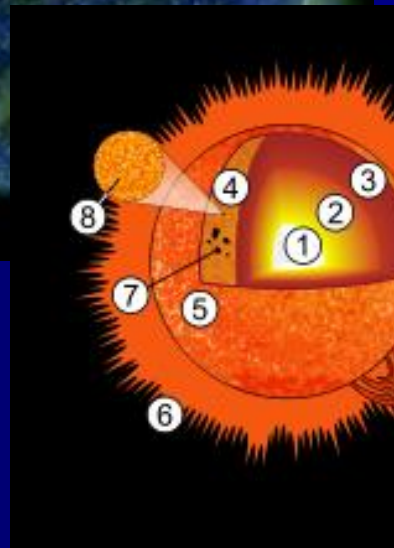
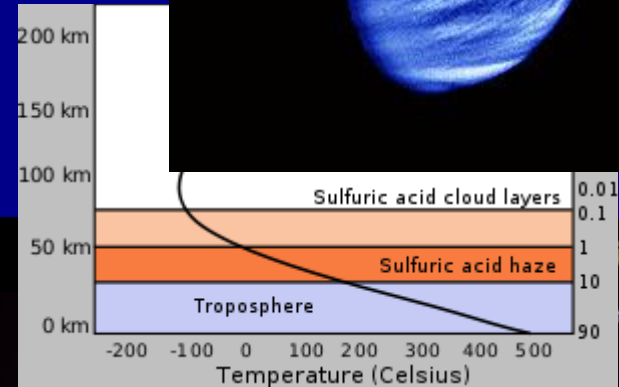
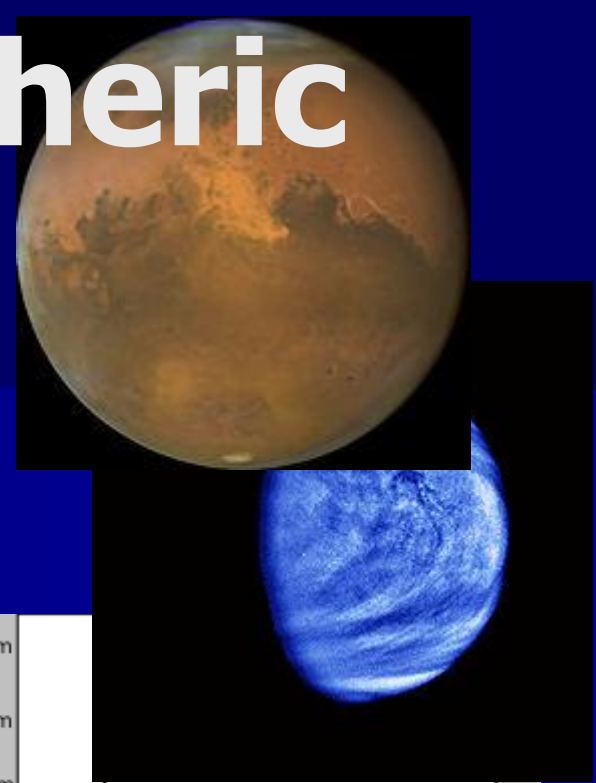
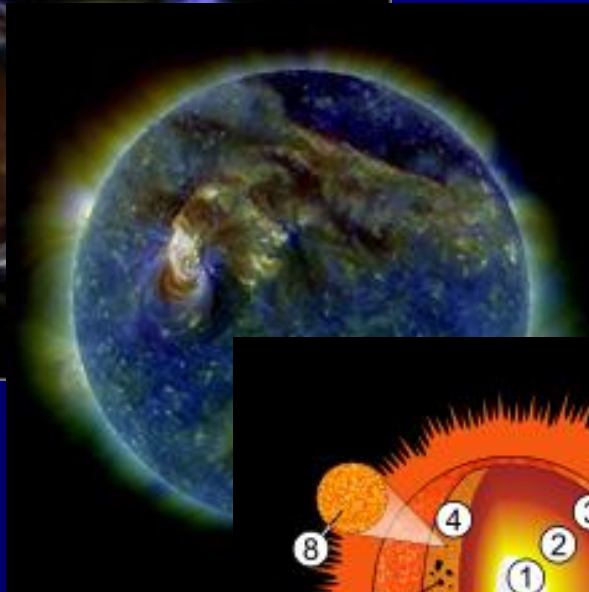
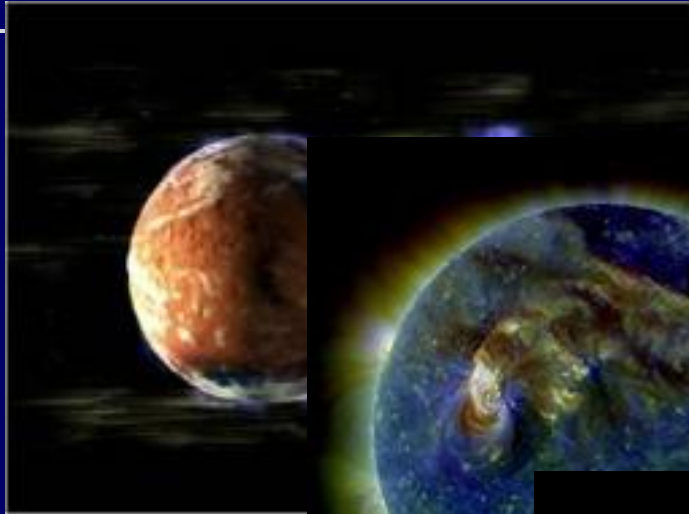


Space &  
Atmospheric  
Physics

# Space & Atmospheric Physics



Lecture – 05

# Earth Atmosphere

Retaining of Gases in the Earth

Major / Minor constituents

Barometric Equation

Scale Height

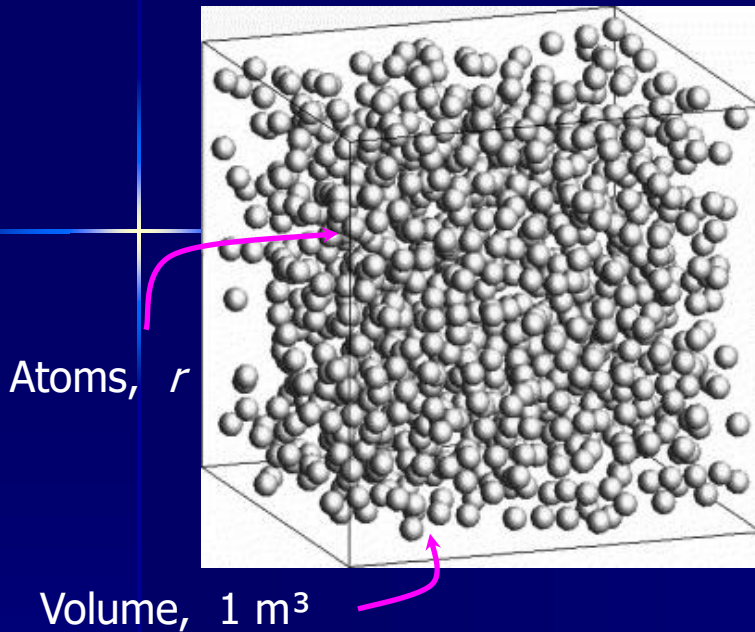
Atmospheric Regions

Temperature Profiles

Retaining of Gases

Number Density Profiles

# Density of the Atoms



**Assume** there are  $r$  atoms in this volume

Masses of the atoms are:

$$m_1, m_2, m_3, \dots, m_r$$

Number densities of those atoms are:

$$N_1, N_2, N_3, \dots, N_r$$

Total Mass of the atoms in the above volume:

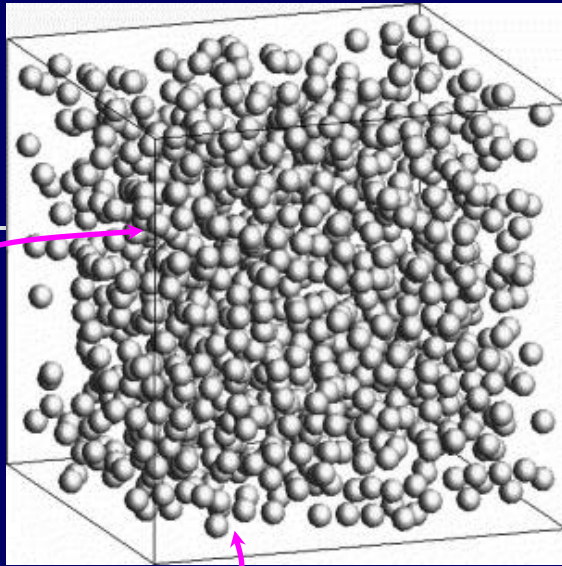
$$m_1 \cdot N_1 + m_2 \cdot N_2 + m_3 \cdot N_3 + \dots + m_r \cdot N_r$$

( This is called the **density** because we consider the unit volume )

Total Molecular Number density:

$$N = N_1 + N_2 + N_3 + \dots + N_r$$

# Density of the Atoms



Atoms,  $r$

Volume,  $1 \text{ m}^3$

Total Mass per  
unit volume

Mean Molecular mass :  $\bar{m}$

$$\bar{m} = \frac{\text{Total Mass}}{\text{Total Molecular Number Density}}$$

$$\bar{m} = \frac{m_1 \cdot N_1 + m_2 \cdot N_2 + m_3 \cdot N_3 + \dots + m_r \cdot N_r}{N_1 + N_2 + N_3 + \dots + N_r}$$

$$\bar{m} = \frac{m_1 \cdot N_1 + m_2 \cdot N_2 + m_3 \cdot N_3 + \dots + m_r \cdot N_r}{N}$$

$$N \cdot \bar{m} = m_1 \cdot N_1 + m_2 \cdot N_2 + m_3 \cdot N_3 + \dots + m_r \cdot N_r$$

Density

# Density of the Atoms

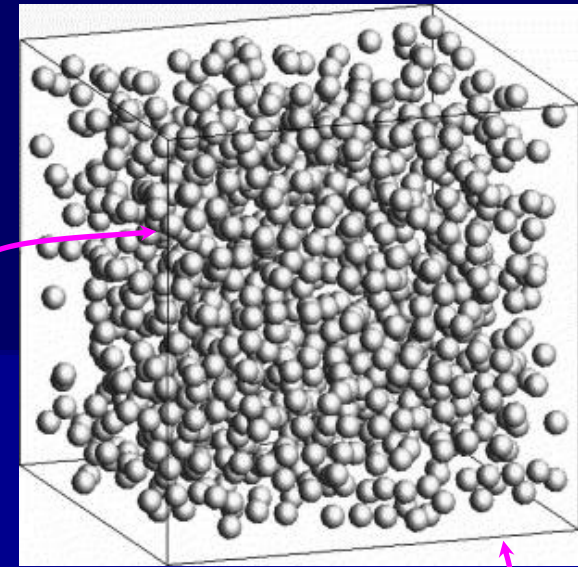
Mean Molecular  
Number Density

Density

$$\rho = N \times \bar{m}$$

Total Molecular Number Density

Atoms,  $r$



Volume,  $1 \text{ m}^3$

# For the Ideal Gas

$$PV = nRT$$

Number of molecules  
per volume,  $V$

$$PV = \frac{NV}{N_o} RT$$

Avogadro Number (Number of  
molecules in a molecular weight)

Boltzmann Constant

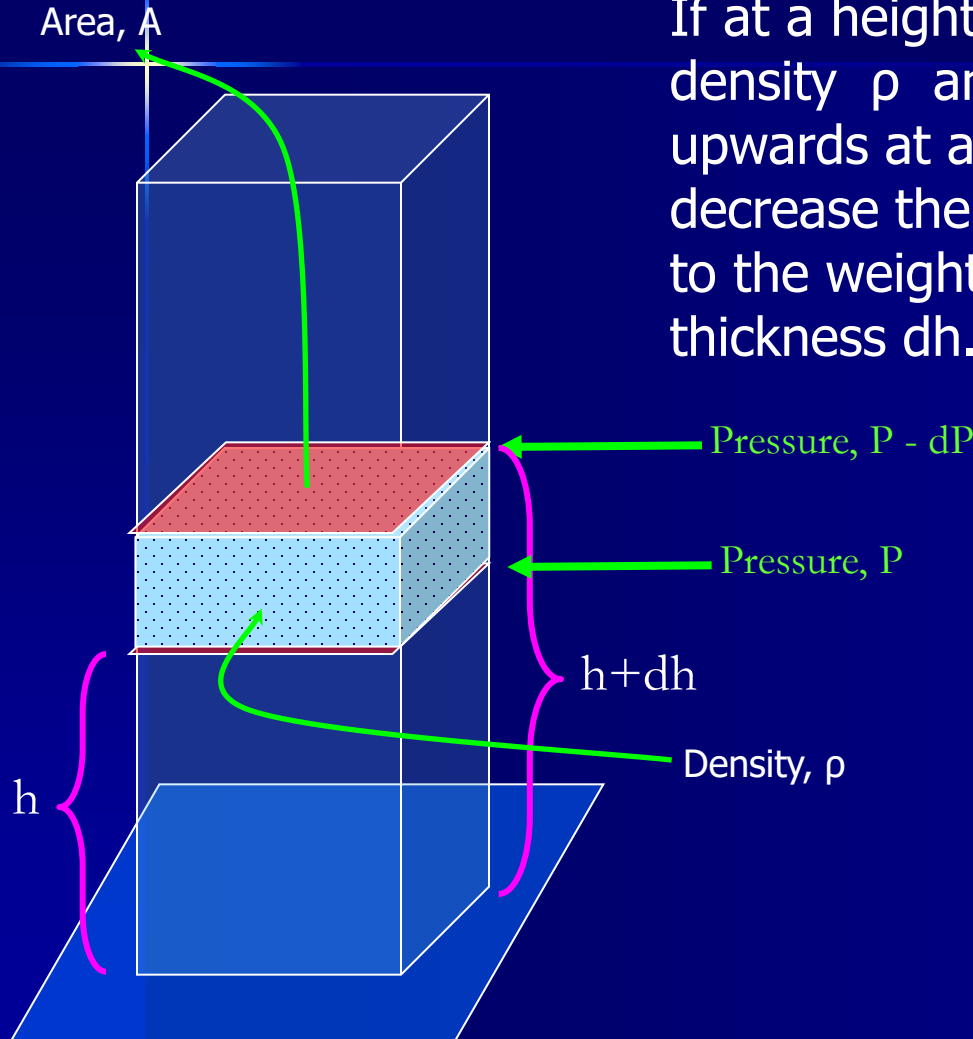
$$P = NkT$$

Where,

$$k = \frac{R}{N_o}$$

# Pressure Profile

The pressure at the Earth's surface (or at higher levels) is a result of the weight of the overlying atmosphere [force per unit area]. If at a height of  $h$  the atmosphere has density  $\rho$  and pressure  $P$  then moving upwards at an infinitesimally small  $dh$  will decrease the pressure by amount  $dP$  equal to the weight of the layer of atmosphere of thickness  $dh$ .



3-D View

$$\begin{aligned} \text{Pressure of the Lower Layer} &= \\ \text{Pressure of the higher Layer} &+ \\ \frac{\text{Weight of the air molecules} \\ \text{in the selected part}}{\text{Cross area of the selected part}} \end{aligned}$$

$$P = P - dP + \frac{A \cdot dh \cdot \rho \cdot g}{A}$$

# Pressure Profile

$$\frac{dP}{P} = -\frac{\bar{m}g}{kT} dh$$

The Pressure at height  $h$  can be written as:

$$P(h) = P_o e^{\frac{-\bar{m}g}{kT} h}$$

This is the general formula as the Pressure at height; This translate as the pressure **decreasing exponentially with height** !

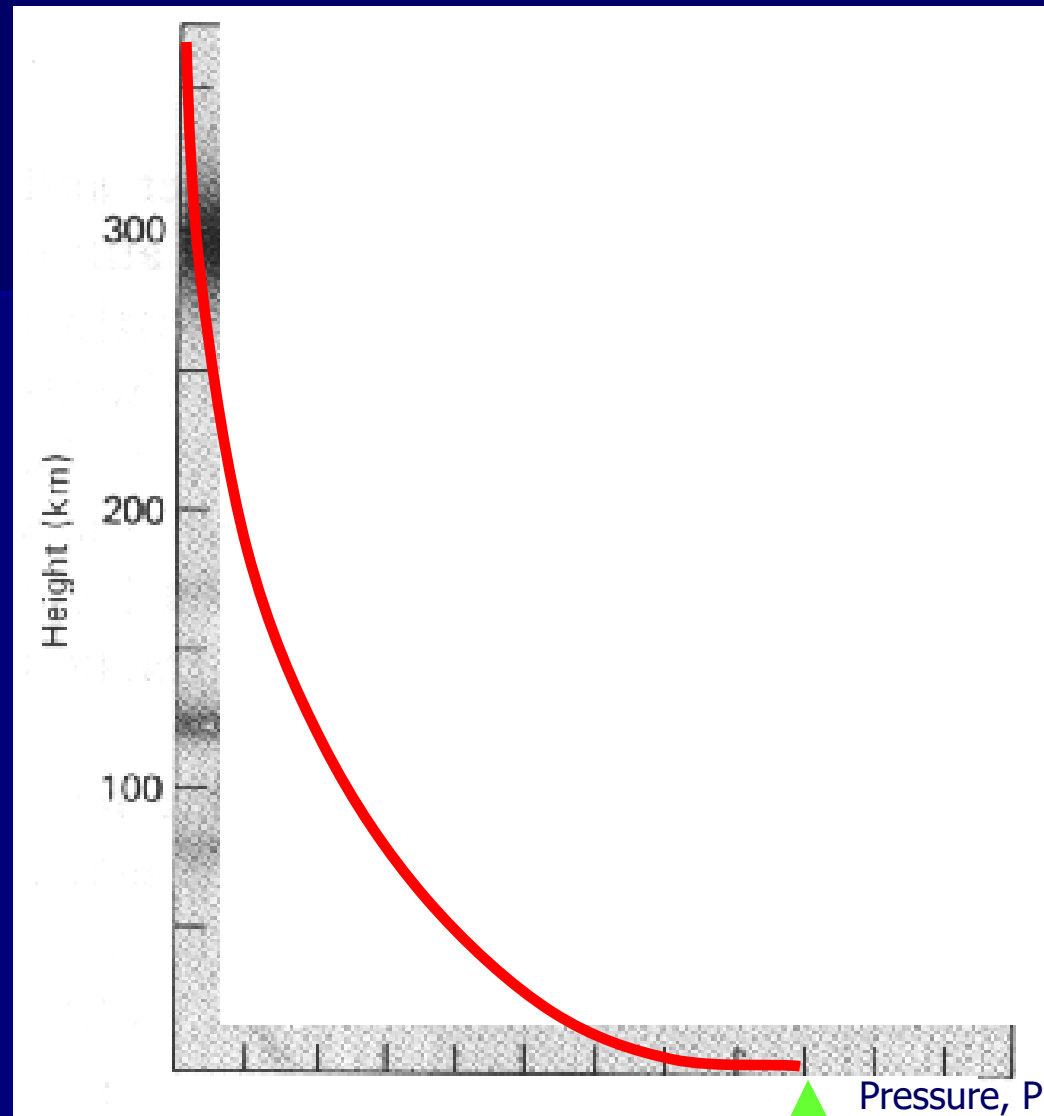
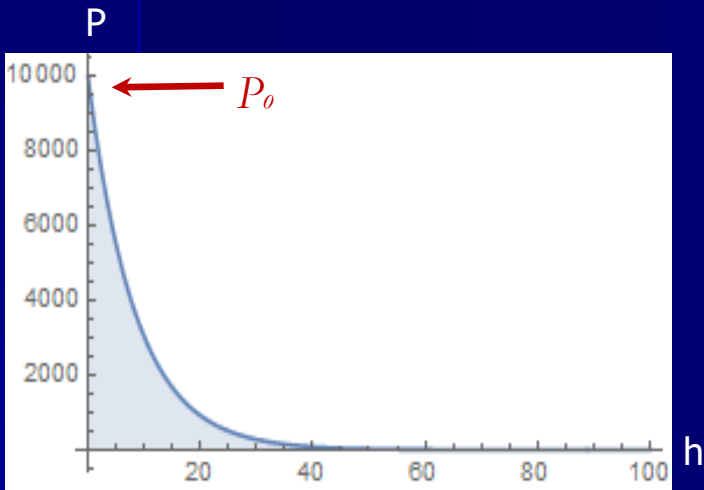
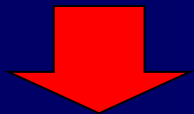
If  $h=0$  then  $P=P_o$  (1); That means  $P_o$  is the pressure at  $h=0$  level or The Ground Level.

Also  $\frac{-\bar{m}g}{kT} h$  is independent of the units. That means  $\frac{kT}{\bar{m}g}$  is also a some height !



# The Graph of $P$ vs $h$ :

$$P(h) = P_0 e^{\frac{-\bar{m}g}{kT} h}$$



# The Graph of $h$ vs $P$ :

$P_0$

# The Graph of $h$ vs $P$ :

## Theoretical Values...

<u>Percent sea level pressure</u>	<u>Altitude (km)</u>
100	0
50	5.6
10	16.2
1	31.2
0.1	48.1
0.01	65.1
0.001	79.2
0.00003	100

Practical Values

```
In[166]:=
po = 100;
m = 5 * 10 ^ (-26);
g = 10;
k = 1.4 * 10 ^ (-23);
t = 300;
ph = 100;
hH = (k * t) / (m * g) /
```

Out[172]= 8.4

Out[173]= {{h -> 0.}}

```
po = 100;
m = 5 * 10 ^ (-26);
g = 10;
k = 1.4 * 10 ^ (-23);
t = 300;
ph = 10;
hH = (k * t) / (m * g) / 100
```

8.4

{{h -> 19.3417}}

```
po = 100;
m = 5 * 10 ^ (-26);
g = 10;
k = 1.4 * 10 ^ (-23);
t = 300;
ph = 50;
hH = (k * t) / (m * g) / 1000
```

8.4

{{h -> 5.82244}}

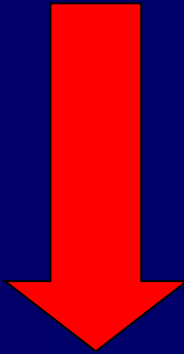
```
po = 100;
m = 5 * 10 ^ (-26);
g = 10;
k = 1.4 * 10 ^ (-23);
t = 300;
ph = 1;
hH = (k * t) / (m * g) / 1000
```

8.4

{{h -> 38.6834}}

# Scale Height (H)

$$H = \frac{kT}{\bar{m}g}$$



where:

- $k$  = Boltzmann constant =  $1.38 \times 10^{-23} \text{ J}\cdot\text{K}^{-1}$
- $T$  = mean planetary surface temperature in kelvins
- $\bar{m}$  = mean molecular mass of dry air (units kg)
- $g$  = acceleration due to gravity on planetary surface ( $\text{m/s}^2$ )

$$H = \frac{(1.4 \times 10^{-23}) \times (300)}{(5.0 \times 10^{-26}) \times (10)}$$



$$H = 8.4 \text{ km}$$

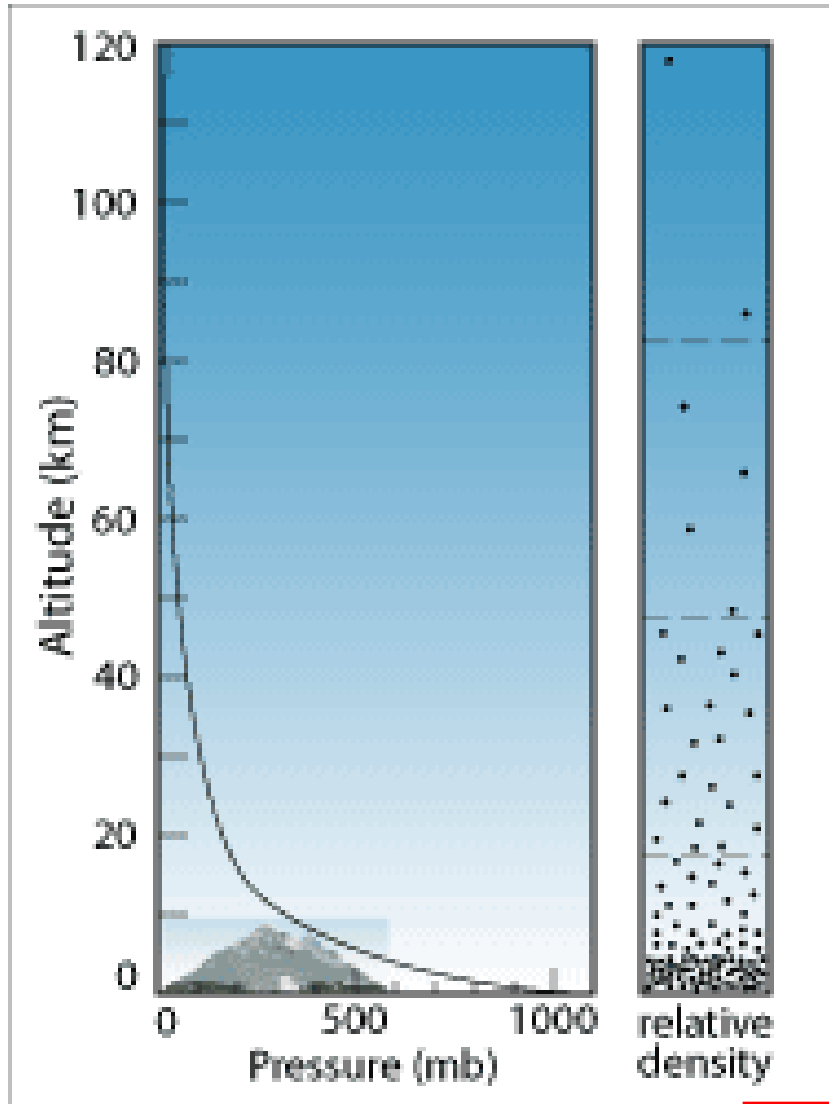
Theoretically this  $H$  is a constant. But practically this  $H$  is not a constant. Because, the values of “**mean molecular mass**”, “**acceleration due to gravity**” and “**mean planetary surface temperature**” are changing with respect to height from the Earth surface.

## The Graph of Scale Heights vs P :

Height	Pressure	
H	$P_o / e$	0.36 $P_o$
2 H	$P_o / e^2$	0.13 $P_o$
3 H	$P_o / e^3$	0.04 $P_o$
4 H	$P_o / e^4$	0.01 $P_o$
5 H	$P_o / e^5$	0.006 $P_o$
.....	.....	
n H	$P_o / e^n$	

## Scale Height of the Earth, $H$

### Temp, $T$ vs Sca Hght, $H$



Pressure and density decrease rapidly with altitude.

T (K)	H (m)
290	8500
273	8000
260	7610
210	6000

bars	millibars	atmospheres	millimeters of mercury
1.013 bar	= 1013 mb	= 1 atm	= 760 mm Hg

Correspondence of atmospheric measurement units.

	Height (km)	Pressure	
6 x 1	6	$P_0 / 2$	$P_0 / 2^1$
6 x 2	12	$P_0 / 4$	$P_0 / 2^2$
6 x 3	18	$P_0 / 8$	$P_0 / 2^3$
6 x 4	24	$P_0 / 16$	$P_0 / 2^4$
6 x 5	30	$P_0 / 32$	$P_0 / 2^5$
	.....	.....	
	6 n	$P_0 / 2^n$	

# Molecular Number Density

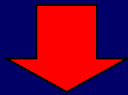
Using the Pressure Equation :

$$P(h) = P_o e^{\frac{-h}{H}}$$

Where,  $H = 8.4\text{km}$

For the Ideal Gas

$$PV = nRT$$



$$P = NkT$$



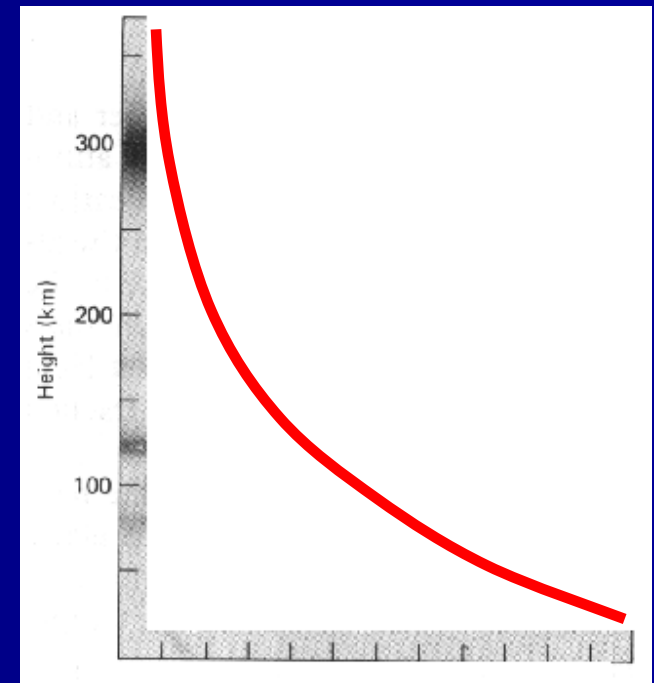
$$N = \frac{P}{kT}$$

$$N(h) = \frac{P(h)}{kT}$$

&

$$N_o = \frac{P_o}{kT}$$

$$N(h) = N_o e^{-\frac{h}{H}}$$



Molecular Number Density

# Molecular Number Density

$$N(h) = N_o e^{-\frac{h}{H}}$$

If  $h = H$ ,

$$N(H) = N_o e^{-\frac{H}{H}}$$

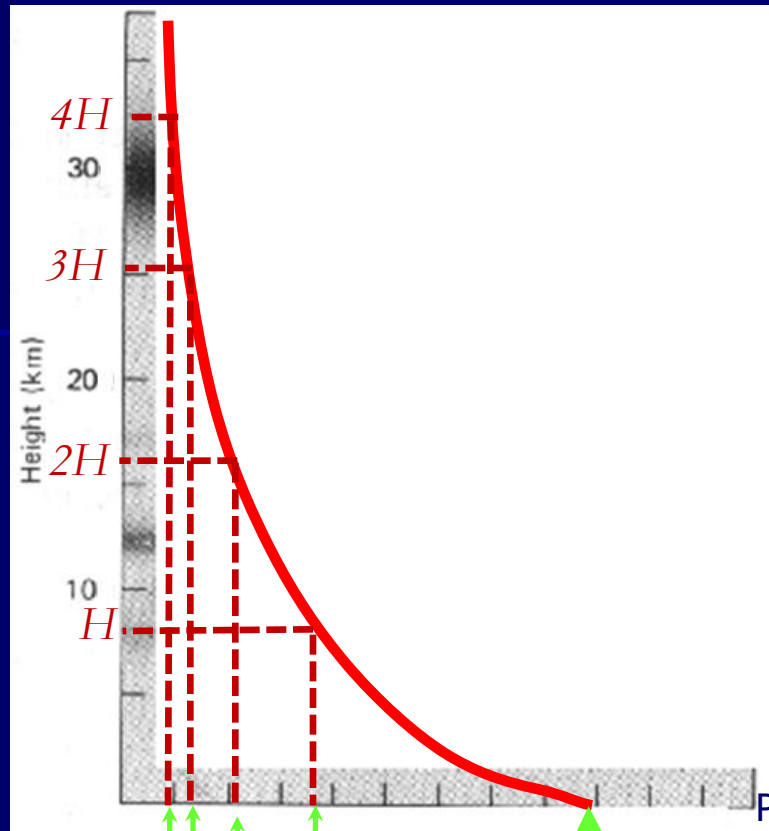
→ 
$$N(H) = \frac{N_o}{e}$$

→ 
$$0.36 N_o$$

Height	Mol Num Den	
H	$N_o / e$	0.36 $N_o$
2 H	$N_o / e^2$	0.13 $N_o$
3 H	$N_o / e^3$	0.04 $N_o$
4 H	$N_o / e^4$	0.01 $N_o$
5 H	$N_o / e^5$	0.006 $N_o$
.....	.....	
$n H$	$N_o / e^n$	



# The Graph of H vs N :



$\frac{N_o}{e}$  36%

$\frac{N_o}{e^2}$  13%

$\frac{N_o}{e^3}$  4%

$\frac{N_o}{e^4}$  1%

$N_o$

**Always Molecular Number Density is decreasing by a factor of e when height is increasing by a multiplies of H**

# Molecular Number Density

$$N(h) = N_o e^{-\frac{h}{H}}$$

Eg:

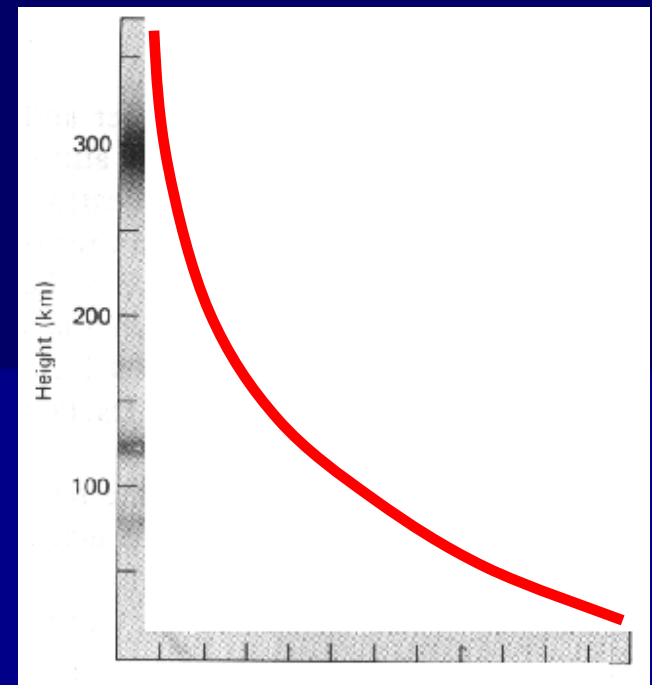
At which height from the surface of the Earth, which you can expect the Molecular Number Density which is half of that of the initial value of the Molecular Number Density ?

If  $N(h) = N_o/2$  when  $h=h$ ,

$$\frac{N_o}{2} = N_o e^{-\frac{h}{H}}$$



$$h \approx 6 \text{ km}$$



Molecular Number Density

	Height (km)	Pressure	
6 x 1	6	$N_o / 2$	$N_o / 2^1$
6 x 2	12	$N_o / 4$	$N_o / 2^2$
6 x 3	18	$N_o / 8$	$N_o / 2^3$
6 x 4	24	$N_o / 16$	$N_o / 2^4$
6 x 5	30	$N_o / 32$	$N_o / 2^5$
	.....	.....	
	6 n	$N_o / 2^n$	

## Molecular Number Density

$$N(h) = N_o e^{-\frac{h}{H}}$$

If  $h=6$  km Then  $N(h) = ?$ ,



$$N = \frac{N_o}{2}$$

If  $h=60$  km Then  $N(h) = ?$ ,



$$N = \frac{N_o}{2^{10}} \approx \frac{N_o}{1000}$$

If  $h=600$  km Then  $N(h) = ?$ ,



$$N = \frac{N_o}{2^{100}} \approx \frac{N_o}{10^{30}}$$

That means at 600 km height, the Molecular Number Density is  $(1/(10^{30}))$  from its initial value.

Consider Linear Distance ;

At 600 km height, the Molecular Linear Distance is  $(1/(10^{30}))^{(1/3)} = (1/(10^{10}))$  from its initial value.

$$= \left( \frac{1}{10^{30}} \right)^{\frac{1}{3}}$$

## Molecular Number Density

$$= \left( \frac{1}{10^{30}} \right)^{\frac{1}{3}}$$

Linear Distance of the molecules = **Mean Free Path** ;  
This is "Separation between two atoms"

Mean Free Path on the ground level =  $6.0 \times 10^{-8}$  m

Mean Free Path at altitude 600 km height from the ground level

:

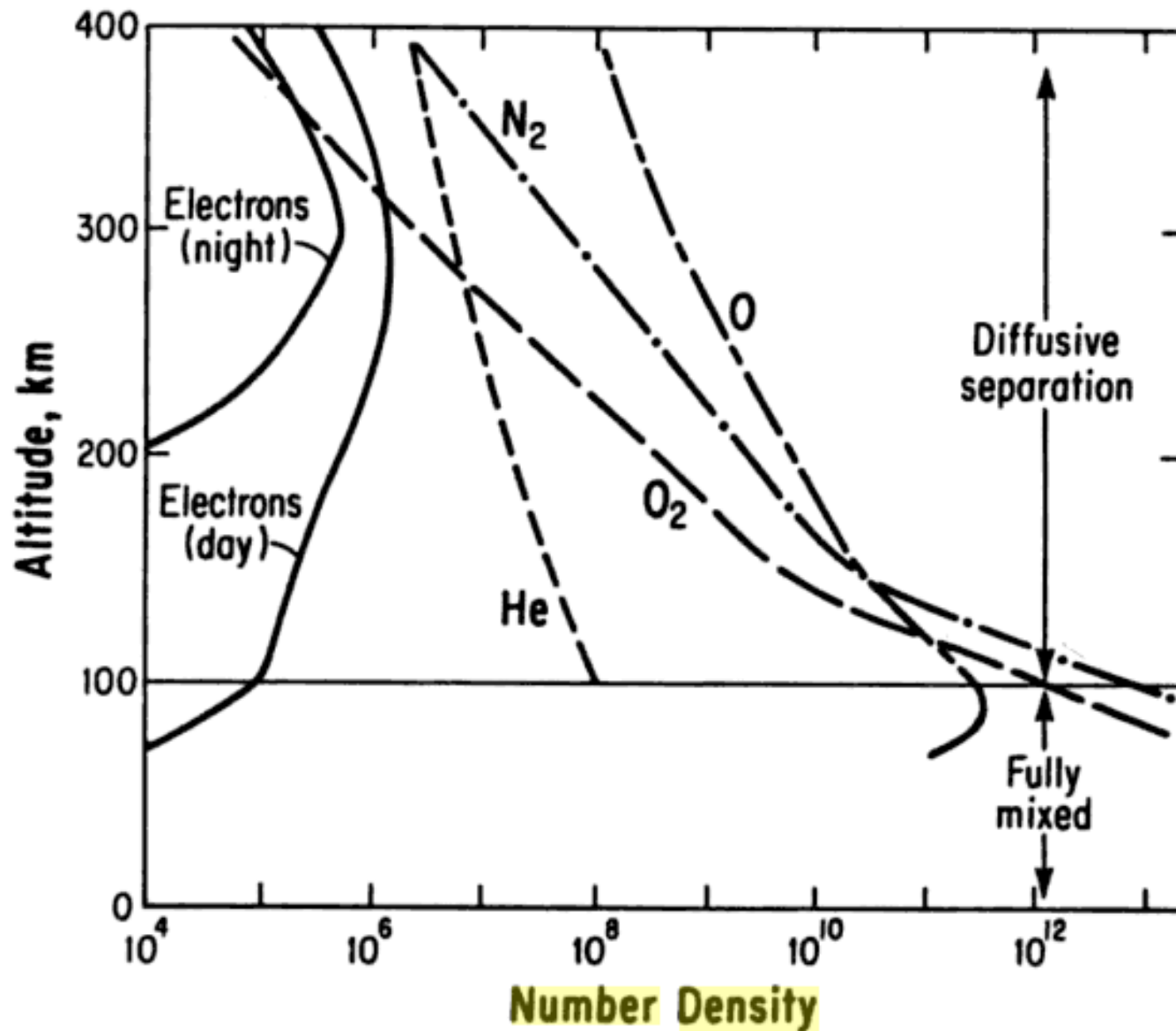
$$= 6 \times 10^{-8} \times (10^{30})^{\frac{1}{3}}$$

$$= 6 \times 10^{-8} \times 10^{10}$$

$$= 600 \text{ m}$$

That means the **gap between two atoms** on that 600 km height (altitude) from the ground level is very high ! At that level there is no mean "The gas", because the **mean free path is very high** (600 m)

# Molecular Number Density



# Density

Using the Molecular Number Density Equation :

$$N(h) = N_o e^{-\frac{h}{H}}$$

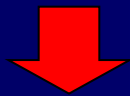
Where,  $H = 8.4\text{km}$

Mean Molecular  
Number Density

Density

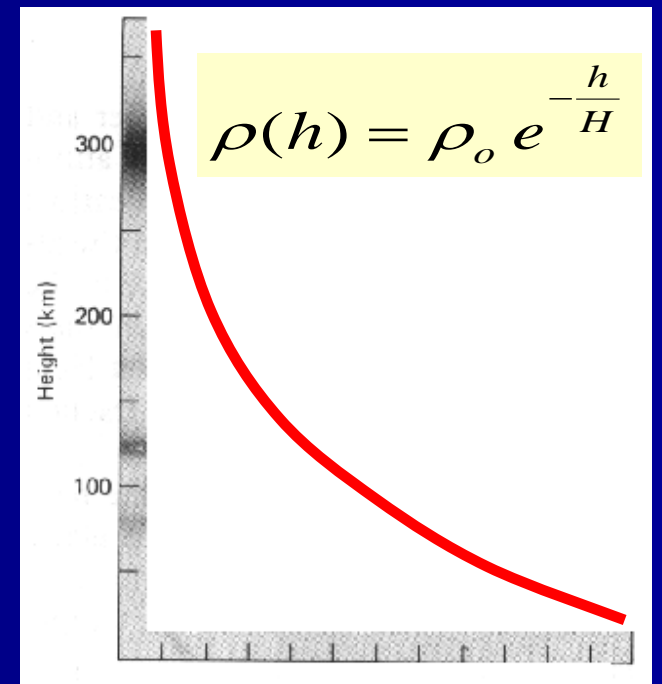
$$\rho = N \times \bar{m}$$

Total Molecular Number Density



$$\rho(h) = N(h) \times \bar{m} \quad \&$$

$$\rho_o = N_o \times \bar{m}$$



Density

# Density

$$\rho(h) = \rho_0 e^{-\frac{h}{H}}$$

Where,  $H = 8.4\text{km}$

If  $h = H$ ,

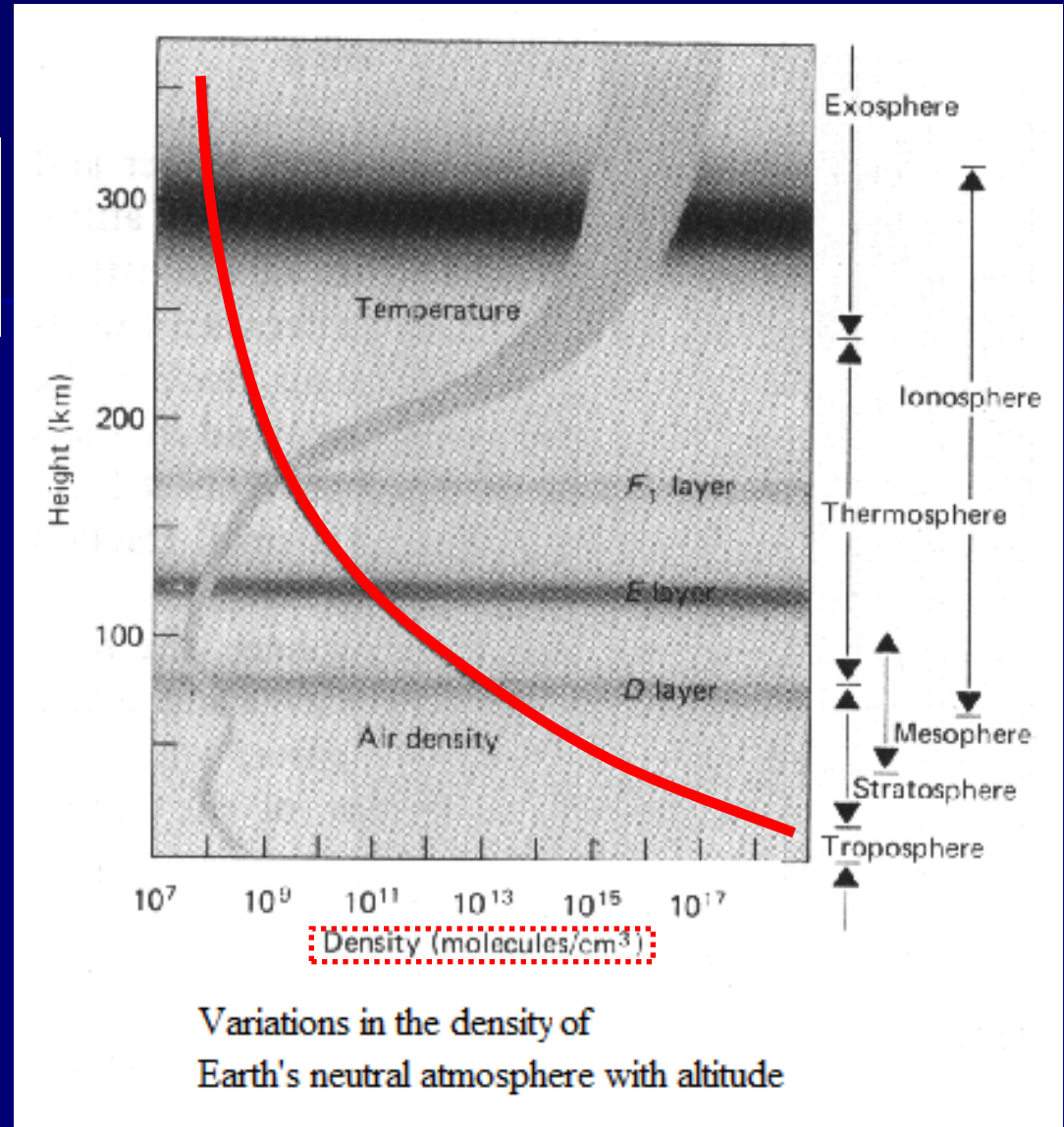
$$\rho(H) = \rho_0 e^{-\frac{H}{H}}$$



$$\rho(H) = \frac{\rho_0}{e}$$

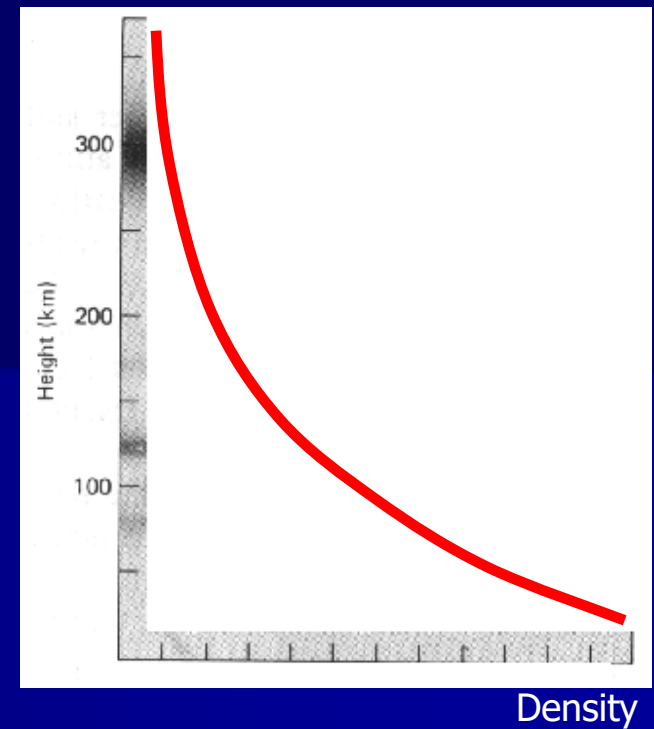


$$0.36\rho_0$$



# Density

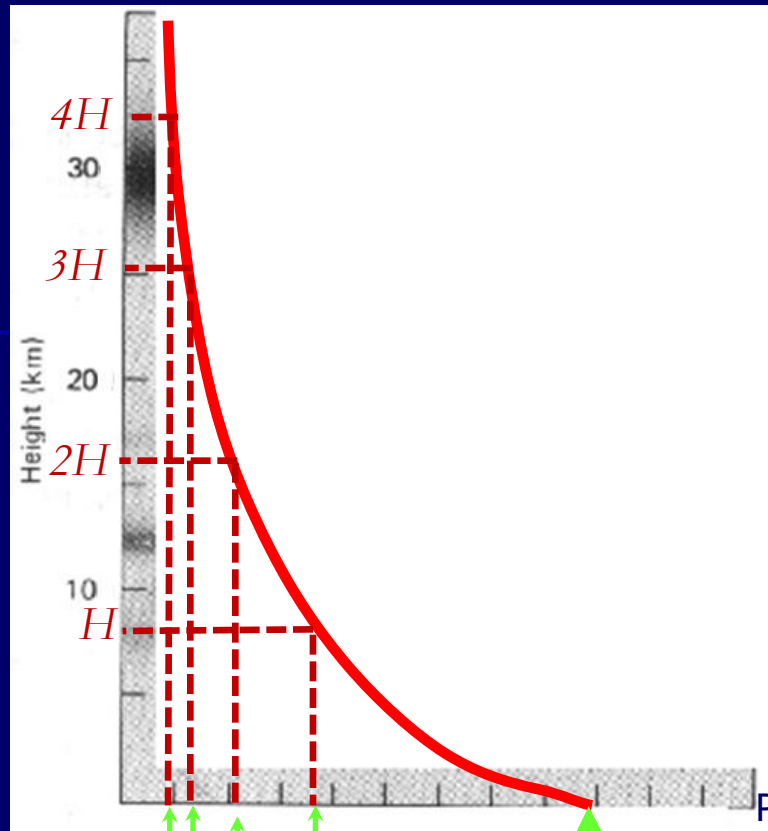
$$\rho(h) = \rho_0 e^{-\frac{h}{H}}$$



Height	Mol Num Den	
H	$\rho_0 / e$	0.36 $\rho_0$
2 H	$\rho_0 / e^2$	0.13 $\rho_0$
3 H	$\rho_0 / e^3$	0.04 $\rho_0$
4 H	$\rho_0 / e^4$	0.01 $\rho_0$
5 H	$\rho_0 / e^5$	0.006 $\rho_0$
.....	.....	
n H	$\rho_0 / e^n$	



# The Graph of H vs $\rho$ :



$\frac{\rho_0}{e}$  36%

$\frac{\rho_0}{e^2}$  13%

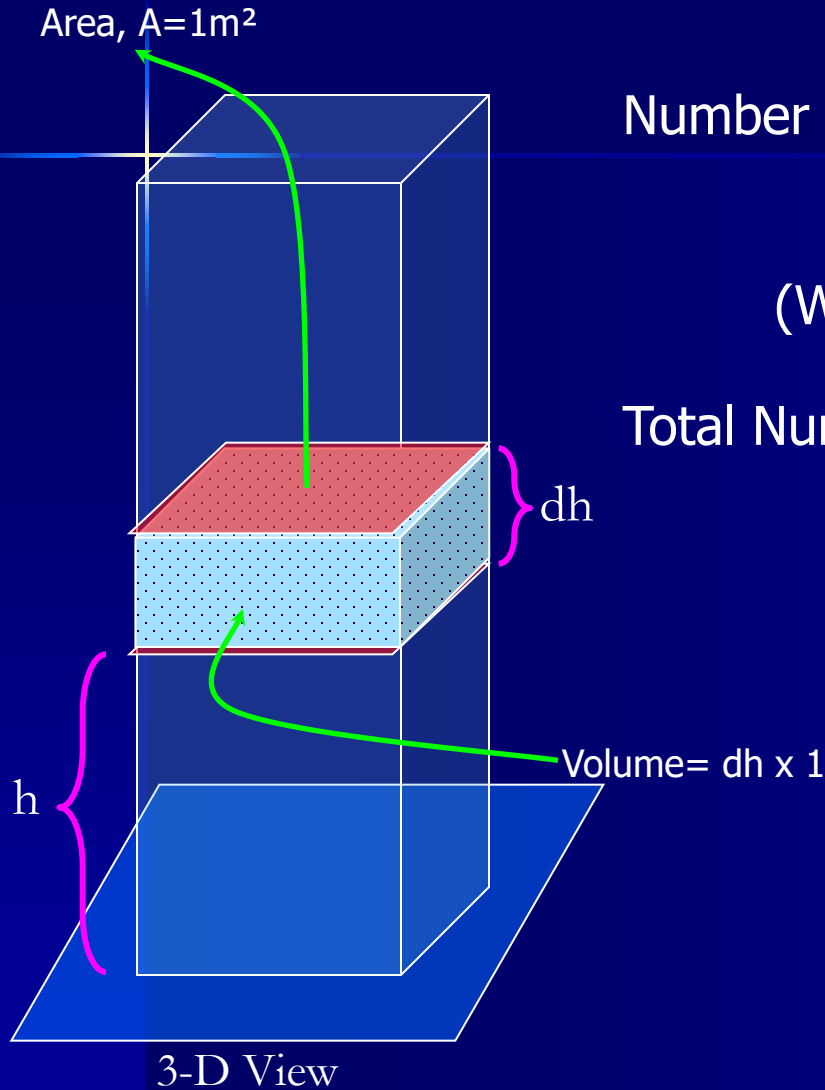
$\frac{\rho_0}{e^3}$  4%

$\frac{\rho_0}{e^4}$  1%

$\rho_0$

**Always Density is decreasing by a factor of  $e$  when height is increasing by a multiplies of  $H$**

# Total Number of Molecules from Earth Surface to altitude $h$ :



Number of molecules in a selected part =

$$N \times dh \times 1$$

(Where  $N$  is the molecular number density)

Total Number of molecules from  $h=h$  to  $h=\infty$

$$\int_{h=h}^{\infty} N \cdot dh$$

Where,

$$N(H) = N_o e^{\frac{-h}{H}}$$

$$\int_{h=h}^{\infty} N_o e^{\frac{-h}{H}} \cdot dh$$

# Total Number of Molecules from Earth Surface to altitude $h$ :

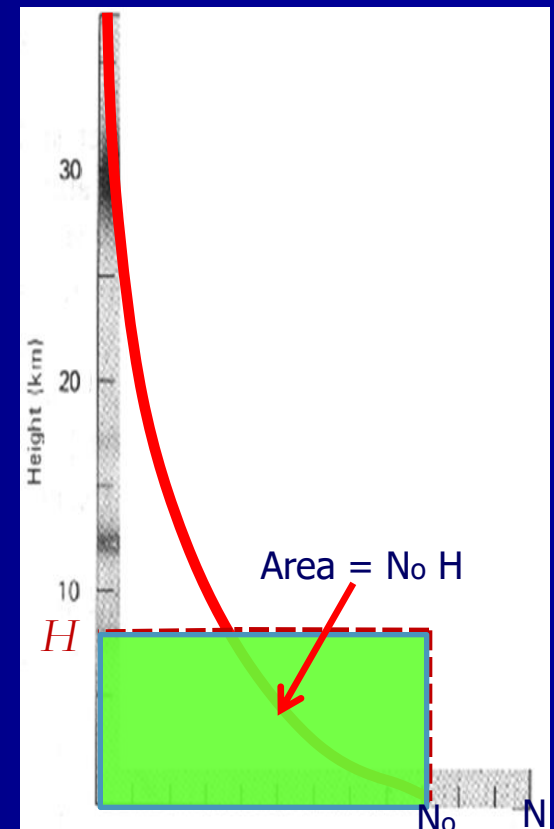
$$N_{Total} = N_o H e^{-\frac{h}{H}}$$

Case I :

$$N_{Total} = N_o H$$

That means, if the molecular number density of the atmosphere of the Earth varies **linearly** without varying **exponentially**, the atmosphere of the Earth will diminish after  $\sim 8.4$  km (a scale height).

This gives to us another definition for the Scale Height !



# Total Number of Molecules from Earth Surface to altitude $h$ :

$$N_{Total} = N_o H e^{-\frac{h}{H}}$$

Case II :

$$\frac{N_{Total}}{N_{Total}} = \frac{N_o H e^{-h/H}}{N_o H} = e^{-h/H}$$

Fraction of the Number of Molecules from the specific height  $h$ .

If  $h=H$  km Then RATIO = ?, 

$$\left( e^{-h/H} \right)_{h \rightarrow H} = e^{-H/H} = e^{-1}$$

$\sim 40 \%$

**60 % of the total molecules exist bellow  $H$  (8.4 km) !**

# Total Number of Molecules from Earth Surface to altitude $h$ :

If  $h=2H$  km Then RATIO = ?, 

$$\left( e^{-h/H} \right)_{h \rightarrow 2H} = e^{-2H/H} = e^{-2}$$

$\sim 15\%$

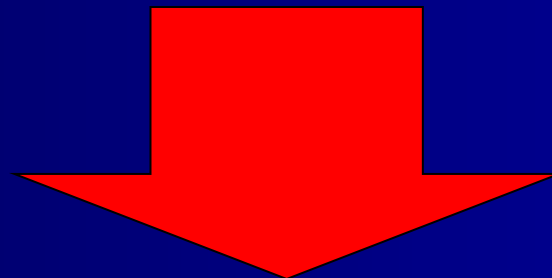
**85 % of the total molecules exist bellow  $2H$  (16.8 km) !**

If  $h=3H$  km Then RATIO = ?, 

$$\left( e^{-h/H} \right)_{h \rightarrow 3H} = e^{-3H/H} = e^{-3}$$

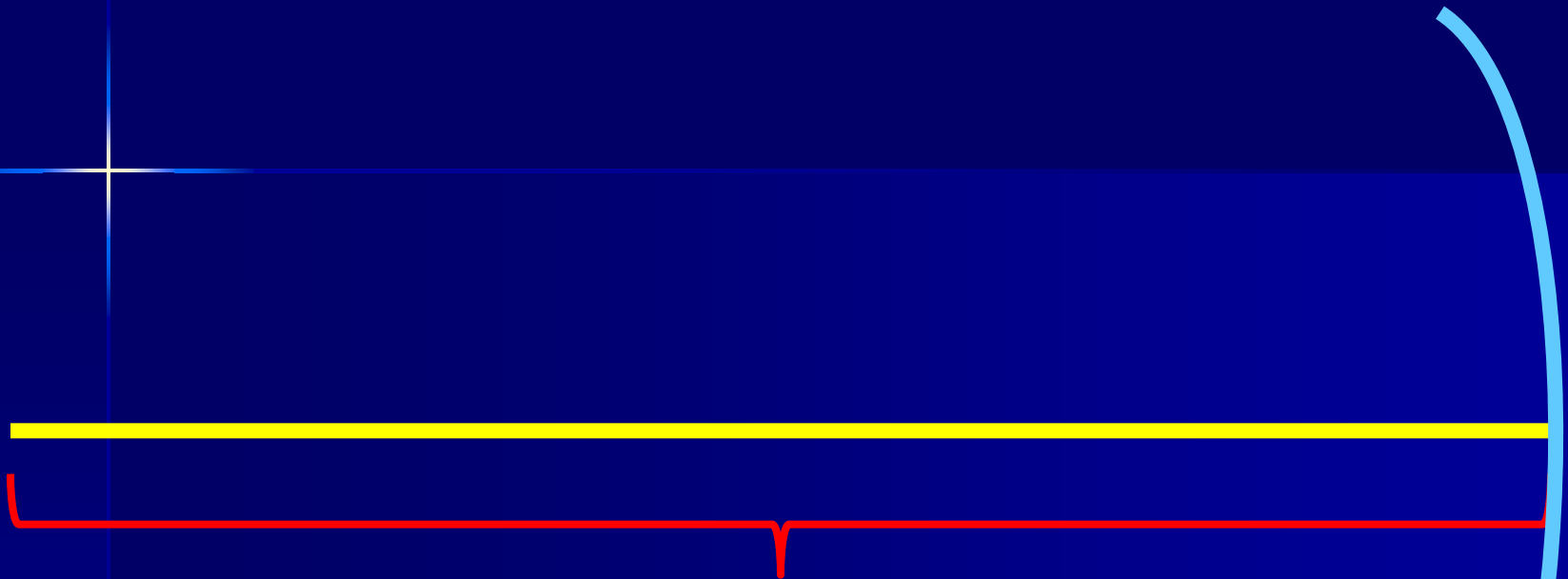
$\sim 5\%$

**95 % of the total molecules exist bellow  $3H$  (16.8 km) !**



h (km)		$N(h \rightarrow \infty) / N(0 \rightarrow \infty)$	% below h
H	08.4	36.78	63.21
2 H	16.8	13.53	86.46
3 H	25.2	4.97	95.02
4 H	36.6	1.83	98.16
5 H	42.0	0.67	99.32
6 H	50.4	0.24	99.75
7 H	58.8	0.09	99.90
8 H	67.2	0.03	99.96
9 H	75.6	0.01	99.98
10 H	84.0	0.004	99.995

# Sketch the size of the Earth's Atmosphere



20 cm straight line

**This is the size of the Earth's Atmosphere**

If we assume the earth to be an Orange which has a radius of 20 cm; then the peel (rind) of the orange is like the atmosphere of the Earth!

1 mm thick line

# Earth Atmosphere

Retaining of Gases in the Earth

Major / Minor constituents

Barometric Equation

Scale Height

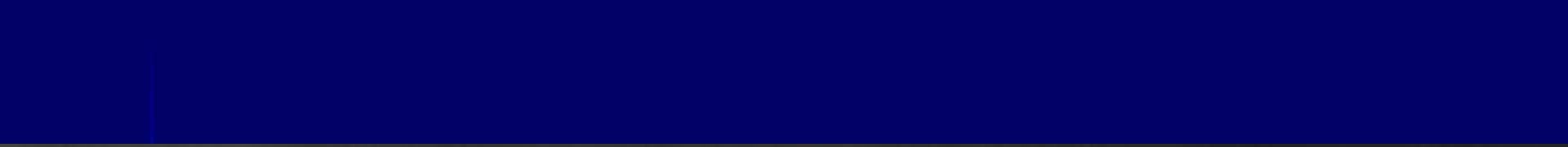
Atmospheric Regions

Temperature Profiles

Retaining of Gases

Number Density Profiles





Thank You !